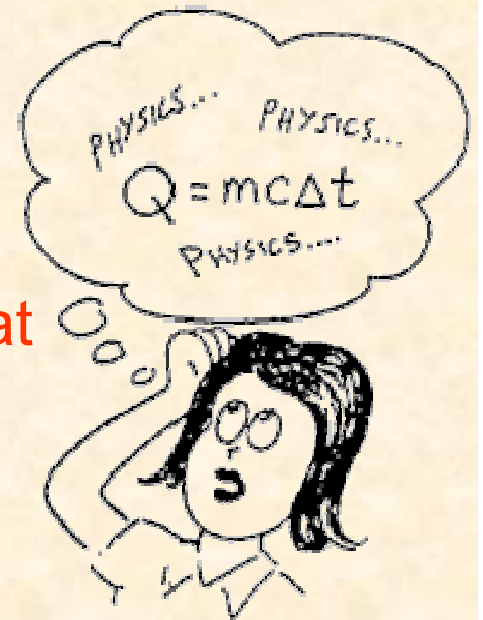


Thermal Histories of Convective Earth Models & Constraints on Radiogenic Heat Production in The Earth

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“Hot” issues

Three main issues of debate mentioned:

- Maintaining surface heat flux for long times – settled with discovery of radioactivity
- Transporting heat from depth to the surface – Conduction? Radiation? Convection?
- **How much heat is now produced?** Can Earth still be warming?

General Constrains

- Earth had to be hot in early stages – considering evidence for core formation and magnetic field, rapid accretion, short-lived radioactive elements.
- Geotherm in the Archean probably steeper- hotter magmas (komatiites) and thinner lithosphere.

Convective Heat Transfer

- Nusselt number - the ratio of total heat flux q through the surface, to the conductive heat q_c , given ΔT :

$$\text{Nu} = qD/K\Delta T = (R_a/R_{ac})^p$$

- Rayleigh number:

$$\text{Ra} = g\alpha D^3 \Delta T/\kappa \nu$$

- Combining them, we get:

$$q/q_0 = (T/T_0)^m$$

$$m = 1 + (n+1)p, \text{ and } p \approx 1/3$$

Large change in q – small change in T

Thermally Activated Rheology

- Deformation by creep of dislocations or by diffusion of vacancies – **thermally activated**
- Diffusion law : $D = D_0 \exp(-H/RT)$
- Effective viscosity: $\nu \sim T \exp(H/RT)$
- $n = - \partial \ln \nu / \partial \ln T = H/RT - 1$
- Observations: mantle viscosity is \sim constant
→ H/RT constant → n is constant, ~ 20 →
 $m = 10$ represents convection

Model Setup

- Thermal state of the interior – represented by a characteristic temperature T
- Considering only post-core segregation
- Assuming a steady-state
- Solving a non-dimensional differential equation:

$$\frac{\partial T}{\partial t} = h - q$$

Non-dim
heat flux

Non-dim heat
production

- Non-dim. time: $t = t' / \tau$, where $\tau = cT_0' / q_0$

Results

- No heat sources

- Exponential decay

- $m=1: T=e^{-t'/\tau}$, $\tau = 15$ b.y.

- $m=10: 1/T^{m-1} = 1+(m-1)t$

- thermal catastrophe 1.5 b.y. ago!

- Constant heat sources:

- $T=q=h+(T_0-h)e^{-t}$

- Approaches a steady state

- Conclusions:

- a planet with constant heat sources will become convective
 - Temperature of convective planet is buffered through thermally activated rheology

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Please see:

Davies, Geoffrey F. "Thermal histories of convective Earth models and constraints on radiogenic heat production in the Earth." *Journal of Geophysical Research* 85, no B5 (May 10, 1980): 2517-2530.

Decaying Heat Source

- Main radioactive elements: U, Th, K.
- Define *decay constants* as $\lambda = \ln 2 \tau / \tau_R$
- Two radioactive abundance models: $\lambda = 6, 9$
- Example results:

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T, h, q as function of time

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q in time, for different h_0 values
transition from "hot" to "cold" past

Constrains On Radioactive Heat Production

- Assuming a **hot** history: q' cannot be ≥ 5 times today's later than 2.5 b.y. ago
→ $h_0 \geq 0.2$
- Assuming a **cold** history: q' cannot be smaller than today's ($q \geq 1$) 1 b.y ago
→ $h_0 \leq 0.7$
- A plausible **warm** history: taking hot temperature for Archean upper mantle → $2 \leq q \leq 5$ 3 b.y. ago → $0.45 \leq h_0 \leq 0.55$ ($\lambda = 9$)

Radioactive Heat Source Concentrations

- Chondritic Coincidence: the heat loss per unit mass of the earth \approx the radiogenic production per unit mass of chondritic meteorites
- Using above values for heat production rate, and observed values for heat loss, conclude:
present heat production rate is very similar to that of carbonaceous or K-depleted chondrites
- Other types of chondrites: rate is too high.

Conclusions

- Imbalance between heat production and heat loss – about half of the heat was generated over the past few billion years
- Heat production to heat loss ratio ≈ 0.5
- Rate of radiogenic heat production similar to carbonateous chondrites
- Results dominated by strong temperature dependence of rheology.